# ESTIMATION OF HYDROLOGICAL SOIL PROPERTIES OF NARSINGPUR DISTRICT



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#### Abstract

This study deals with the estimation of hydrological soil parameters of the doab between Sher and Barau river and Bargi left bank canal falling in Narsingpur district of M.P..The soil properties to be determined are Soil Texture, Sat. Hydraulic conductivity using Guelph permeameter (for in situ measurement) and from Johnsons graph, Matric flux potential, Alpha and Soil moisture characteristic curves.

Based upon the soil classification of Soil Survey Department of Govt. of M. P., a base map for soil classification was prepared and at twelve locations disturbed soil sampling and insitu measurement of Saturated conductivity was performed. The soil samples were brought to the Soil Water Laboratory of the Institute and the textural analysis of all the samples was carried out and is presented here. The matric flux potential, alpha parameter and saturated hydraulic conductivity as given by Johnsons were also determined and are presented here.

#### 1.0 INTRODUCTION

water and air, which compete for the same position in soil in the root zone, are both needed for plant growth. The soil moisture deficiency is abated through irrigation and the oxygen deficiency is done away with by providing drainage facilities. If the primary object is to lower the ground water table below a certain level below root zone of the crops, then sub surface drainage system is recommended to be installed.

In sub surface drainage design texture is an important parameter which governs the movement of water in side the soil. The design and functioning of sub surface drainage systems depends to a great extent on soil's saturated hydraulic conductivity. All sub surface drain spacing equations make use of this parameter.

Plants require oxygen as well as water for their growth. During waterlogged condition oxygen diffusion is unable to sustain root or microbial requirements of plants for any length of time. In the absence of sufficient oxygen, substance such as alcohol and cyanide may be formed in the plant tissues and plant growth is curtailed. The allowable groundwater table is closely related to the root system of the crop concerned.

The present area under study is the doab between Sher and Bareau river. It falls under Narsingpur district in the Jabalpur division of Madhya pradesh. City Narsingpur is about 100 kms away from Jabalpur city. This area comes under the command area of Bargi irrigation project, a multipurpose project. The Bargi dam is constructed on river Narmada. The area for the present study selected is the doab of river Sher, Baraue and left bank canal of Bargi irrigation project.

# 1.1 Salient features of the Bargi Project:

The proposed Bargi Project in M.P. comprises of a dam about one mile long and 210' high across the river Narmada at Bargi near Jabalpur, and canal system to provide controlled irrigation to 6.6 lac. acres, a major portion of the area lies towards the south west and the remaining to the north-east of Narmada river. About 20,653 acres is under forest while the remaining area is under cultivation.

Whereas, the river discharge touches a peak point of about 4.19 lac cusecs during the monsoon season, it falls off very considerably during the summer months. Its catchment area up to the proposed dam site comprises of 5,600 square miles of partly hilly and partly forest covered land. The gross capacity of the proposed reservoir is 3.18 million acre feet of which 0.60 million acre feet will be allocated for dead storage.

It is proposed to take off canal system which would irrigate land on both sides of the river. It will irrigate area of 6,60,000 acres on both the banks of the river,

#### Location:

1. State Madhya Pradesh

District Dam - Jabalpur

Command - Jabalpur

& Narsinghpur District

3. Latitude 22 - 56' - 30")

Longitude 79 - 55' - 30" }ToposheetNo.55.

4. River Narmada

Location Dam site is near village

Bijora 11.2 km (7 miles) east of village Bargi

# Hydrology:

		C.G.S. (units)	F.P.S. (Units)
1.	Catchment area upto dam site.	14,556.05 sq.km.	5,620 sq.M.
2.	Maximum annual rain- fall 1926 upto (Jamtara)		<b>90.98</b> inches
3.	Minimum annual rain- · fall 1899 (upto Jamtara)	664 mm.	26.16 inches
4.	Average annual rain- fall 1891-1967 (upto Jamtara)	1,448 mm.	57.01 inches
5.	Design flood	45,296 m <sup>3</sup> /sec.	16,00,000 cusecs
6.	Actual observed maximum flood, at Jamtara (9-7-61)	11,876 m <sup>3</sup> /sec.	4,19,500 cusecs
7.	Available runoff at Bargi (1891-1967) from computer series		
i 	i) 50% dependability ii) 75% dependability iii) 90% dependability	0.672 mham. 0.471 mham.	7.250 M.A.F. 5.450 M.A.F. 3.880 M.A.F.
	l System (Irrigation):		
	Gross command area		
2.	Culturable area	2.98 lakh hac.	7.36 lakh acre
3.	Area to be irrigated	2.67 lakh hac.	6.60 lakh acre

4. Annual irrigation 5.181 lakh hac. 12.80 lakh acre

<b></b>	Left bank main canal	107.58 M <sup>3</sup> /sec.	3,800 cusec		
	Right bank main canal	58.06	2,051 "		
	Right bank main canal	37.86 "	1,337 "		
Leng	Length of Main Canal:				
i.	Left bank main canal	135.06 km.	83.94 miles		
ii.	Right bank main canal No. I	43.65 km.	27.13 miles		
iii.	Right bank main canal	95.57 km.	59.40 miles		

# 1.2 General description of the area:

## 1.2.1 Location & extent:

The area lies between E.longitudes 78 - 35' to 79 - 50' and N. latitudes 22 - 35' to 23 - 5' and includes Jabalpur district and Gadarawara and Narsinghpur tahsils of Narsinghpur district. In the north, river Hiran limits the boundary of command area of Jabalpur district, while river Narmada limits the boundary for Narsingpur district. In the west, the extent of command area ends with Shakker river. In the south and west direction there is no prominent features to show the boundary of the area. However, proposed left bank canal will form the boundary in south. The area to be irrigated tahsilwise is as follows:

Tab 1. Tahsilwise area to be irrigated from Bargi Project in Narsingpur distrct (acres)

District	Tehsil	Geographical area as per villagewise statistics	Cultivable area	
			Covered by Project	To be irrigated
Narsinghpur	Narsinghpur	3,05,653	2,45,743	2,21,166
	Gadarwara	1,24,680	1,01,144	91,033
	Total	4,30,333	3,46,887	3,12,199

# 1.2.2 Physiography & relief:

The elevation above mean sea level of the command area varies for Narsinghpur from 1027 to 1248 feet. The general topography of the area appears to be flat except in the vicinity of the rivers, where deep gullies and ravines have formed giving rise to undulating to rolling topography. As such the entire area is a broad plain of low relief, Local differences in elevation is small due to adoptions of Haveli system of cultivation, which has checked the erosion. In the plain area, the slope ranges from 0 to 3%, but in area having undulating topography the steeper slopes even up to 15% are noticed. The distribution of area under different slope class are as per table 1.

Table 2: Distribution of area under different class

Slope class	Slope %	Area covered	Percentage
A	0 - 1 %	2,56,620	29.65
В	1 - 3 %	2,68,480	31.03

		8,65,232	100.01
CD	3 - 8 %	87,532	10.11
вс	1 - 5 X	64,000	7.39
E	8 - 15 %	17,920	2.07
D	5 - 8 %	10,240	1.18
С	3 - 5 %	1,60,440	18.58

# 1.2.3 Rivers, major streams & drainage:

The area is provided with a number of rivers and The main river of the area is Narmada, it covers of the area starting from village Ghughri in Jabalpur village shakalpur in Gadarwara tahsil. The depth of Narmada in various localities varies from 40 to 100 feet. The river Narmada flows from south to north from village Ghughuri to Gwarighat than onward, it flows east to west. The other major tributaries are Sher and Sakko both flow south-east to north-west and merge with Narmada at Sagon Ghat and Sokalpur respectively. number of minor tributaries namely Bururewa, Umar, Machhavai in Narsinghpur Singri, Sitarewa and Shanker nala, In this way district is having a net work of drainage, which drain excess water of the area successfully and finally into the river Narmada. On an average after every 10 miles one or the other major or minor streams are present for the drainage. streams and rivers are quite deep and do not silt up like those in alluvium tracts, where the natural drains have chocked up and have blocked the drainage system. Hence due to net work of drainage and their position in the landscape the command area is well drained. The natural drainage follows the general slope of the land. The rainfall of Narsingpur varies from 45" to 50", but even there is no problem of water stagnation or upward movement of salts.

#### 1.2.4 Geology:

Soils of whole of the Bargi Project in Jabalpur and Narsingpur district have been derived from trap rook but with regard to soils only Deccan trap is important one which has given rise to Characteristic colour and properties to the soils of the area. Deccan trap is the great formation of horizontally bedded basaltic lavas that occupies a large portion of the western part of India.

In Narsingpur district quite a large area of the district is covered by the Deccan trap. Fresh trappen exposures are seen in the south-east and north of the district while the central part is occupied by the older alluvium and flat land composed of dark grayish brown, dark brown with small area under yellowish brown colour.

#### 1.2.5 Climate:

The tract enjoys a sub-tropical climate. The annual rainfall of Narsinghpur district varies from 45 to 50" most of it precipitate during monsoon season i.e. end of June to end of September. The summer temperature goes as high as 45 C. The extremes of cold and heat are experienced during winter and summer respectively. The average data of different years are presented in table 3.

- 2.0 Methodology
- 2.1 Particle Size Distribution

#### 2.1.1 Soil Texture

Soil texture is a term which refers to the size range the particles present in the soil mass. The diameter particles present in the soil sample makes the soil to be coarse, medium and fine. Table 3 gives the textural class names material larger than 2 mm in diameter. The traditional method of characterising particle sizes in the soils is to devide the particle sizes into gravel, sand, silt and clay. The soil is actually determined by separating these fractions and measuring their proportion which is called the mechanical analysis. The soil texture triangle is then used to convert quantitative data from detailed gradation analysis of separates less than 2 in diameter to textural class names of soils. Soil texture ài especially important in sub surface drainage as it has a direct relationship with hydraulic conductivity and water retention (David 1982).

Tab 3: Textural classification as per particle Diameter

Material	Diameter
Stones	>10 inches
Cobbles	10 - 3 inches
Coarse gravel	3 - 0.5 inches
Fine gravel	0.5 - 2 mm
Very coarse sand	2 - 1 mm
Coarse sand	1 - 0.5 mm
Medium sand	0.5 - 0.25 mm
Fine sand	0.25 - 0.1 mm
Very fine sand	0.1 - 0.05  mm
Silt	0.005 - 0.002 mm
Clay	<0.002 mm

In a particle size distribution curve, the y-axis or the ordinate in the graph indicates the percentage of soil particles having the diameter finer than indicated on X-axis. Chart are also available in literature showing the percentages of clay, silt and sand in the basic soil textural classes. The triangle of particle size distribution is given in Fig 2 (drainage manual USDA).

#### 2.1.2 Soil Moisture Tension and PF values

The moisture contained in the pore spaces of a soil mass is subjected to the capillary forces. This capillary force causes a negative soil moisture tension which is also called suction. The suction is expressed as the height of water column (h) that will rise from the water table against the force of gravity. This height is inversely proportional to the diameter of the pores. Therefore, h = 0.3/d, where d = equivalent pore diameter (EPD) of a cylindrical pore with the same capillary. The negative logarithm of soil moisture tension in centimeters of water is used to indicate the soil moisture tension. This negative log of soil moisture is referred to as pf. Tab below shows the characteristic figures of capillary soil moisture.

Tab 4: Characteristic figures of soil moisture and pF

Equivalent pore	SMT	pF value
diameter ( m)	(cm)	
3,000	1	0
1,200	2.5	0.4
1,000	3.0	0.5
300	10	1.0
30	100	2.0
20	150	2.20 FC
15	200	2.30
9	340	2.51 ME

3	1,000	3.00
0.3	10,000	4.00
0.2	15,000	4.18 WP
0.03	10 <sup>5</sup>	5.00
0.003	10 <sup>7</sup>	7.00 oven dry

# 2.1.3 Field Capacity

The water present in a saturated soil is allowed to drain out, the water quickly leaves the soil via largest pores and air is pulled into the soil. This movement of water is mainly due to the gravitational potential difference. When the rapidly moving water in the unsaturated soil ceases to move then the soil is said to be at Field capacity. Field capacity occurs when soil the maximum amount of water with little or no further loss of water by drainage or loss of gravitational water. soi1 Α matrix potential of about -1/3 bars has been found to correspond to the field capacity. A bar is equal to 10 dynes. A dyne is equal to the force that imparts to a mass of 1 gram an acceleration of 1 cm/sec. A bar is also equal to 1,020 cm water column or 1,020  $gm/cm^2$ .

# 2.1.4 Wilting Point

As soil becomes drier, the conductivity rapidly decreases and movement and uptake of water becomes slower. Therefore, if no additional water is added to the soil, the plant will absorb water slower than water is lost by transpiration. Thus a water deficit develops in the plant. This point is called wiling point. A soil water matrix potential of about -15 bars has been found to correspond to wilt point (Henry, 1984).

#### 2.1.5 Available Water

The water present in the soil between field capacity and wilting point is known as available water. It is generally considered to be matrix potential in the range of -0.3 to -15 bars.

#### 2.1.6 Effect of texture on Available water

The capacity of soil to hold water is related to surface area as well as pore space volume. Hence, water holding capacity is related to both structure and texture of the soil. In fine textured soils have the maximum total water holding capacity, but maximum available water is held in medium-textured Several researchers have indicated that available water in many soils is closely correlated with content of silt and very fine sand. It is well known that sandy soils are more droughty than clayey soils, because fine-textured soils are able to retain available water. Also, there is a difference in the soil of soil-moisture characteristic curves of sand and clay. The flatness of the curve for fine sandy loam at water matrix potential is less than -4 bars which means that most of the available water sandy soils have a high potential. Therefore, plants can use this water in sandy soils. Since in clay or clay loam the water is available at lower potential therefore it can be rapidly used by the plants.

#### 2.3 Sieve Analysis

In the Indian Standard (IS: 460-1962), the sieves are designated by the size of aperture in mm, whereas in BS (410-1962) and ASTM (E11-1961) standards, the sieve sizes are given in terms

of the number of openings per inch. These are described in Seth, 1990. The mechanical analysis of soil is performed by sieve analysis and sedimentation analysis. The sieve analysis is basically carried out in two parts i.e. coarse analysis and fine analysis.

The soil sample for which the sieve analysis is to carried out is first dried in the oven. The dried sample is then sieved through 4.75 mm sieve (Indian Standard). The retained on the sieve is known as the gravel portion. The portion which passes through the 4.75 mm sieve is used for finer sieve analysis. The sieves used for fine sieve analysis are: 2mm, 425, 300. 212, 150 and 75 micron is sieves. In order to carry out the sieve analysis the sieves are arranged in one over another in the lowering order of mash openings. The cover is placed on top of assembly and a container is kept at bottom. assembly is shaken by sieve shaker. Shaking of the sample performed for nearly 10 minutes. The soil sample which is to analysed and which is passing through 4.75 mm sieve is washed with distilled water. The washing is done to dislodge the silt clay particles sticking on sand particles. For providing dispersion of different particles two grams of andium hexametaphosphate is added per litre of water used. Washing is done till water coming out through 75 micron sieve is perfectly clean. The portion retained on the 75 micron sieve is then dried in oven. The sample is then placed on the top sieve of sieve assembly and sieve is shaked.

At the end the portion retained on each sieve is collected and weighted. The percentage of soil sample retained on each sieve on the basis of total weight of soil sample and the percentage of weight passing through each sieve was calculated (Bowles ,1986).

The calculations were started with 100 percent and subtracting the percentage retained on each sieve as a cumulative procedure as given by

Percentage passing = Percentage arriving - percentage retained

2.4 Wet Mechanical analysis

The soil fraction which is finer than 75 micron size is used for sedimentation analysis. This analysis is based on the stokes law which states that all other factors being constant the velocity at which grains settle out of suspension is dependent upon the weight, shape and size of grains. Assumptions are made for the analysis that the soil particles are spherical and all the particles have some specific gravity. This assumption leads to the fact that coarser particles settle more quickly than finer ones. The terminal velocity of a particle in suspension is given by following formula:

$$v = -\frac{D^2}{18} \frac{\gamma_s - \gamma_w}{\eta} \qquad (1)$$

where, v = Terminal velocity,
r = Radius of spherical particle (cm),
D = Diameter of particle (cm),
y = density (unit wt.) of particles (g/cm³),
y = density of water/liquid (g/cm³)

n = viscosity of water/liquid (g-sec/cm²)
y = µ/g
p = viscosity in absolute units of dynes-sec/sq.cm,
and g = acceleration due to gravity (cm/sec)

If water is used as a medium for suspension,  $\gamma_s = G \gamma_w = G$  as  $\gamma = 1$  g/cu.cm. Therefore from eq. (1)

$$v = \frac{G-1}{1800\eta} p^2$$
 .....(2)

At  $20^{\circ}$  c viscosity of distilled water is approximately 0.01 poise. For G = 2.68, the equation (2) reduces to

$$v = 91.5 \text{ p}^2$$
....(3)

Equation (3) is an approximate solution of stokes law and is used for estimation of diameter of soil particles (Seth, 1990).

# 2.5 Hydrometer Method

In the hydrometer analysis the weight Wd per ml of suspension is found by reading the density of soil suspension at a depth He at various time intervals. This height He goes on increasing as the particles settle with the increase in time interval. Therefore, it is necessary to calibrate the hydrometer and sedimentation jar before the start of the sedimentation test. This calibration will provide the relation between He and the density readings of the hydrometer.

# 2.5.1 Calibration of Hydrometer

The stem of the hydrometer has horizontal markings which gives the density of the soil suspension situated at the centre of the bulb at any time. For the sake of convenience the hydrometer readings are subtracted by one and the remainder is multiplied by one thousand to give a reduced reading indicated as Rh. Hydrometer readings increase in the downward direction towards hydrometer bulb. Referring when hydrometer is immersed in the jar water level increases, the level as rises to alal and bb to bibl. At this point time bibl corresponds to the centre of hydrometer at which density measurements are taken.

He = (H + h/2 + Vh/2A) - Vh/AHe = H + 0.5(h - Vh/A)

Above equation have two variable He and H which depends upon the hydrometer reading (Rh). By selecting various hydrometer readings the depth H can be measured with the help of an scale and corresponding effective depths (He) can be found as Vh, A and h are constant for a given hydrometer.

# 2.5.2 Test procedure

# 2.5.2.1 Soil Suspension Preparation

About 24 to 60 gm of oven dried sample (depending upon the type of soil) is taken and is weighed accurately. The sample placed in a beaker and distilled water is added to form a smooth thin paste. A deflocculating agent (e.g. sodium oxalate, sodium silicate and sodium polyphosphate compound such as tetra muiboa petrophosphate, sodium hexametaphosphate (calgon) and sodium tripolyphosphate) is also added to get a proper dispersion of IS:2720 (Part IV)-1965 recommends the use of dispersing soil. solution containing 33 g of sodium hexametaphosphate and 7g of sodium carbonate in distilled water to make one litre of solution. 50 ml of this solution is added to beaker containing soil this mixture is warmed gently for about 10 minutes. The contents are then transferred to the cup of mechanical mixture. The soil sample is stirred well for about 15 minutes and is then washed through 75 micron sieve using distilled water. Now the sample ready for analysis and is transferred to measuring cylinder care should be taken that the volume of suspension should bе 1000 m 1 for analysis.

The sedimentation jar is shaken vigorously and is kept

verticle over solid bases. The hydrometer is inserted and readings are taken at definite time interval (1/2, 1, 2, 4, 8, 5, 30 mins and 1,2,4 hours etc.). To take reading, hydrometer is inserted 30 seconds before the given time interval so that it is stable at the time when reading is taken. The reading is taken for upper level of meniscus. A suitable meniscus correction is then applied to the hydrometer readings.

#### 2.5.2.2 Correction for hydrometer reading

Hydrometers are generally calibrated at 27°C, if temperature of soil suspension is not 27°C. a temperature correction (Ct) should be applied to the observed hydrometer If temperature is more than 27°C the reading of hydrometer will be less hence temperature correction will be positive and vice versa. Other corrections to be applied are meniscus correction and dispersing agent correction. As reading of hydrometer is taken at the top of meniscus, actual reading at water level is higher. Hence meniscus correction (Cm) is: positive. It is found by immersing the hydrometer in clean The dispersing agent correction (Cd) is always negative as it increases the density of water.

Therefore corrected hydrometer reading

R = Rn + Cm + Ct - Cd

where, Rn = observed hydrometer reading at the top of meniscus

Cm, Ct and Cd can be combined into C which is called composite correction.

R = Rn + C

# 2.5.2.3 Determination of composite correction

To calculate 'C' an identical cylinder with 1000 ml capacity is taken and filled with distilled water and same quantity of dispersing agent as is used in sedimentation analysis. The temperature of both cylinders being same the hydrometer is immersed in this comparison cylinder. The reading of hydrometer is taken at the top of meniscus. The negative of this reading so obtained gives the composite correction. Composite correction is found before the test and also at time interval more than 30 minutes.

## 2.5.2.4. Computation of D and N

The observed hydrometer readings were corrected suitably as per the calculated meniscus and temperature corrections and from the corrected hydrometer reading the He was calculated for the corresponding time and Rh. The particle size diameter was calculated by

$$D = 10^{-5} M (He/t)^{0.5} \dots (4)$$

The over all percentage finer was calculated as below

$$N = N' - \frac{W1}{W} - \dots$$
 (5)

where, N' = 
$$\frac{100 \text{ G}}{\text{Wd}} = \frac{100 \text{ G}}{\text{G}} = \frac{1}{1}$$

N = over all percentage finer,

N' = percentage finer based on Wd,

G = specific gravity of soil particles,

Wd = weight per ml of suspension, and

W = total dry weight of soil

t = time to fall through a height He

The grain size was plotted against percentage finer on semi log paper and are given in figs 3 to 6 for all the samples.

2.6 Measurement of Field Saturated Hydraulic Conductivity by Guelph Permeameter

Hydraulic conductivity is the measure of the ability of a soil to conduct water under a unit hydraulic gradient. saturated hydraulic conductivity (Ks) refer to the saturated hydraulic conductivity of soil containing entrapped air. saturated hydraulic conductivity is more appropriate than the truely saturated hydraulic conductivity for unsaturated zone investigations because by definition, positive pressure not persist in unsaturated conditions long enough for entrapped air to dissolve. In the presence of the water table, the hole method is a simple and reliable technique for measuring saturated hydraulic conductivity in relatively uniform However this method cannot be used if the water table is present in the region of interest. The methods for measuring hydraulic conductivity in the absence of the water table are more complicated. The shallow well permeameter method, also known as the dry auger hole method and the bore hole permeameter are the techniques for measuring hydraulic conductivity. Hydraulic conductivity decreases as the soil water suction increases. This called is the conductivity pressure head relationship relationship. The Guelph permeameter is used to determine Ks for a particular soil. Once the soil water suction is measured, hydraulic conductivity (K) for that soil at that soil water suction  $(\varphi)$  can be readily estimated by relatioship

Guelph permeameter can measure matric flux potential  $(\phi_m)$  which is the measure of a soil's ability to pull water by capillary force through a unit cross sectional area in a unit time. The matric flux potential  $(\phi_m)$  in sq.cm/sec is given by following relationship

$$\phi_m = 0.0572 * X * R1 - 0.0237 * X * R2.....(7)$$

Alpha parameter  $(\alpha)$  is the slope of the curve relating the natural log of hydraulic conductivity (K) to soil water suction in per cm expressed by following relationship

$$\alpha = K_2 / \phi_m$$
 ....(8)

# 2.6.1 Guelph Permeameter Apparatus

The Guelph Permeameter is essentially an "in hole" Mariotte bottle constructed of concentric transparent plastic tubes. The apparatus comprises the following sections. The two models of Guelph Permeameter can be reffered in Shukla and Soni (1993).

- (i) Tripod Assembly
- (ii) Support Tubes and lower air tube fittings
- (iii) Reservoir Assembly
- (iv) Well Head Scale and upper air tube fittings
- (v) Auxiliary tools

#### (i) Tripod Assembly:

The tripod assembly consists of a tripod based with movable tripod bushing and three detachable tripod legs complete with end tips. The flexible tripod base has three leg sockets into which the tripod legs are inserted. Tripod chain is used for firm

placement and support of tripod legs.

# (ii) Support Tube and Lower Air tube fittings:

These are the fittings which conduct water from the reservoir assembly into the well hole and provide the for establishing and maintaining a constant head in the well hole. The support tube supports the reservoir assembly over the well and transports water from the reservoir to the water outlet. water outlet tip serves as a base for the permeameter disperses the energy of the out flowing water through the ribbed vents at the bottom to the tip to minimize erosion of the soil in the well hole. The air tip seating washer rests on the inside step of the water outlet tip and is the seat for the Air Inlet tip when against the air tip seating the air inlet is fully seated washer, air cannot move up through the support tube and there no flow of water out of the reservoir. The air inlet tip is connected to the bottom of the lower air tube and is used regulate the well head height. The air restriction washer located inside the air inlet tip and regulates air flow to provide a constant, non fluctuating head in the well.

# (iii) Reservoir Assembly:

The reservoir assembly provides a means of storing water and measuring the outflow rate while the Guelph Permeameter is in use. It consists of inner reservoir tube, outer reservoir tube, reservoir valve, base and reservoir cap. For studies in very low permeability soils, for example clay soil, use of the inner reservoir alone is required to provide adequate outflow rate. When working in moderate to high permeable soils, for example sands and loamy soils, the reservoir combination is used. The

inner reservoir tube is graduated in centimeters for measuring the rate of fall of water out of the reservoir in both situations. The Guelph permeameter shows the closed or sealed state with air inlet tip sealed against air tip seating washer. When air tube is uplifted, with accompanying air inlet tip and well height indicator, water flows from the reservoir down the inside of the support tube through the water outlet tip and into the well. The water height in the well is established by the height of the air inlet tip. This water height in the well can be set read using well height indicator in conjunction with the well head scale.

The reservoir base includes the reservoir valve. The base connects and seats the inner and outer reservoir tubes to the support tube. Water flow is controlled by the position of the reservoir valve. When the valve position is up, both reservoirs supply water to the well hole. When it is pointing straight down, only the inner reservoir supplies water to the well The reservoir cap provides an airtight cover for the top of the reservoir, the seal of the air tube and supports the well scale. The middle air tube is located inside the inner reservoir tube. Two ports are located in the reservoir cap namely Fill port and Fill plug. The vacuum port consists of an Access Neoprene tube and clamping ring. The vacuum port facilitates pulling a vacuum when the reservoirs are not initially completely filled.

# (iv) Well Head Scale and Upper Air Tube Fittings:

The upper air tube is connected to be Middle air tube with an air tube coupling. It serves as an extension to facilitate setting the well head after the well head scale is put in place.

# (v) Auxiliary Tools

The Guelph permeameter kit includes a soil auger for excavating a well, a sizing auger, a well prep brush, a vacuum hand pump for pulling a vacuum in the reservoir and a collapsible water container for carrying water to the field. The well prep brush meant for removing any smear layer that exists in the augered well hole that may create a barrier to the natural flow of water out of the well into surrounding soil.

# 2.6.2 Procedure

The Guelph permeameter method (Reynold et.al. measures the steady state liquid recharge necessary to maintain depth of liquid in an uncased cylindrical well a constant finished above the water table. Constant head level in the well hole is established and maintained by regulating the level of the bottom of the air tube which is located in the centre of permeameter. As the water level in the reservoir falls, a vacuum is created in the air space above water. When the permeameter operating, an equilibrium is established. The reduced pressure in the air above the water in the reservoir together with the pressure of the water column extending from the surface of well to the surface of the water in the reservoir which is always equal to the atmospheric pressure.

When a constant well height of water is established in a cored hole in a soil, a bulb of saturated soil with specific dimension is rather quickly established. The bulb is very stable and its shape depends on the type of soil, the radius of the well and the head of water in the well. The shape of the bulb is numerically described by the C factor used in the calculations.

Once the bulb shape is established, the outflow of water from the well reaches a steady state flow rate which can be measured. The rate of this constant outflow of water, together with the diameter of the well and height of water in the well can be used to determine the field saturated hydraulic conductivity of the soil.

The Richard analysis of steady state discharge from a cylindrical well in unsaturated soil, as measured by the Gue 1ph permeameter technique accounts for all the forces that contribute to three dimensional flow of water into soils, the hydraulic push of water into soil, the gravitational pull of liquid out bottom of the well and the capillary pull of water out of the well into the surrounding soil. The Richard analysis is the basis for the calculation of field saturated hydraulic conductivity. The C factor is a numerically derived shape f :tor which is dependent on the well radius and head of water i well (Shukla and Soni, 1993).

#### 2.6.3 Procedures for Field Use

Before making a measurement with the Guelph permeameter in the field, it is necessary to perform a site and soil evaluation, prepare a well hole, assemble the permeameter, fill the reservoirs, and place the permeameter in the well hole.

# 2.6.4 Well preparation

The instruments needed for excavating and preparing a well bore hole are soil auger and sizing auger. The soil auger and sizing auger. The soil auger is used to remove bulk amounts of soil and rock. The sizing auger is used as a finishing tool to produce a proper sized well hole of uniform geometry and to

clean debris off the bottom of the well hole. The sizing auger is designed to produce a hole that is uniformly 6 cm in diameter with a flat bottom. Generally, the procedure is to use the soil auger to excavate the well hole down to a depth 15 cm less than that desired for the final well hole. The last 15 cm can than be excavated using the sizing auger to produce a debris free well hole of uniform geometry.

In the moist soils or in medium to fine textured soils, the process of augering a hole may create a smear layer which can block the natural flow of water out of the well into the surrounding soil. In order to obtain reliable and representative results using the Guelph Permeameter, the smear layer must be removed. The well prep brush is designed to use in the standard 6 cm diameter well hole.

#### 2.6.5 Permeameter Placement

Tripod is centered over the well hole and slowly the permeameter is lowered so that the support tube enters the into well hole. The tripod is used to support the permeameter in depth. For use in wells well down to approximately 38 cm in deeper that 38 cm, the tripod bushing alone provides the functions After centering and stabilizing the permeameter. the permeameter is placed, it can be easily filled with water. The following standard procedure should be followed for making measurements.

- (i) Verify that both the reservoirs are connected. The reservoirs are connected when the notch on the reservoir valve is pointing up.
- (ii) Establish a 5 cm well Head Height (H1). Slowly raise the air inlet tip to establish the 5 cm well head height.

Raising the air tube too quickly can cause turbulence and erosion in the well.

- iii) Observe the rate of fall of the water level in the reservoir. If it is too slow, then turn the reservoir valve so that the notch is pointing down. Water will then be supplied, only from the small diameter inner reservoir which will result in a much greater drop in water level between readings.
- (iv) Measure permeameter outflow. This is indicated by the of fall of water in the reservoir. Readings should be made at regular time intervals, usually 2 minute intervals The difference of readings at consecutive interval used. divided by the time interval equals the rate of fall water, R1 in the reservoir. Continue monitoring the rate of fall of water in the reservoir until the rate of fall does significantly change in three consecutive time This rate is called R1 and is defined as intervals. the "Steady state rate of fall" of water in the reservoir at height H1 which is the first well height established and is always 5cm in the standardized procedure.
- (v) Establish 10 cm Well head height (H2). Slowly raise the air inlet tip to establish the second well head height of 10 cm. Monitor the rate of fall of water, R2, in the reservoir until a stable value of R2 is measured.
- (vi) The field saturated hydraulic conductivity, Ks can be calculated using the following equation:

Ks = 0.0041 X R2 - 0.0054 X R1 .....(9) where,

X = Reservoir constant, equal to 35.39 where reservoir

combination is used and 2.14 when only inner reservoir is used

- R2 = Steady rate of fall of water in the reservoir for a head of 10 cm.
- R1 = Steady rate of fall of water in the reservoir for a well head of 5 cm.

### 2.6.6 Pressure Plate Apparatus

This is a standard method for obtaining the soil moisture retention curve. Pressure plate apparatus consists of a pressure chamber in which a saturated soil sample is placed on a porous ceramic plate through which the soil solution passes but no particle or air can pass. The soil solution which passes through the membrane is in contact with atmospheric pressure. As soon air pressure inside the chambers are raised above atmospheric it takes excess water from the soil out of the chamber through the membrane outlet. Soil water will flow out from the soil sample until the metric potential of the unsaturated flow is The air pressure is then same as the applied air pressure. released and the moisture content of the soil is gravimetrically determined.

During a run, soil moisture will flow from around from each of the soil particle and out through the ceramic plate until such time as the effective curvature of the water film through out the soil are the same as at the pores in the plates. When this occurs an equilibrium is reached and the flow of moisture ceases. When air pressure in the chamber is increased, flow of water from the samples starts again and continue until a new equilibrium is reached. A source of regulated gas pressure is required for all extraction work. Compressed air from a compressor is the most

efficient source of supply.

The ceramic plates are available in different range. Each ceramic pressure plate cell consists of a porous ceramic covered on one side by a thin neoprene diaphragm sealed to edges of the ceramic plate. An internal screen between the plate and diaphragm provides a passage for flow of water. An outlet stem running through the plates connects this passage to outflow tube fitting which to the atmosphere outside of extractor. Cross section view of ceramic pressure plate cell soil sample is shown in Fig. 2. To use the ceramic pressure plate cell, one or more soil samples are placed on the porous ceramic surface held in place by retaining rings of appropriate height. The soil samples together with the porous ceramic plate are saturated with water. This is usually done by allowing an excess of water to stand on the surface of the cell for several When the saturation is complete, the cell can be mounted into pressure vessel. Air pressure is used to effect extraction moisture from the soil samples under controlled conditions. The 1 bar ceramic plates are ideal for the routine determination of 1/1 bar and 1/3 bar range of the soil suction. The 3 bar pressure The 15 bar ceramic plate cells are used in the range of 0-3 bars. cells are commonly used for measurement of soil moisture suction in the range of 5-15 bars of soil suction.

The moisture retention curve of a soil sample can generally be determined by equilibrating a soil sample at a succession is known tension value and each time determining the amount of moisture. The graph is plotted between the tension and corresponding soil moisture value to obtain the soil moisture retention curve. Different types of soil yields different retention curves.

#### 3.0 Procedure

#### 3.1 Soil Sampling

The pilot area as shown in fig 1 was selected for the estimation of soil hydrological properties of the area. The soil sampling was carried out from this area. Samples have been collected from 8 different locations as indicated in the fig 1. Measurement of field saturated hydraulic conductivity was also carried out in these locations for both the layers.

# 3.2 Grain Size Analysis

- Oven dried sample weighing 500 gm was taken and soaked with water.
- This sample was washed through no. 200 sieve. The washing was carried out carefully using distilled water.
- Two group of soils, one passing through the sieve and another retained on the sieve were collected separately.
- 4. Both the groups of soil were then oven dried. The group retained on sieve was subjected to sieve analysis and the group passing through the sieve was subjected to sedimentation analysis.

#### A. Sieve Analysis

- Oven dried soil sample retained on no.200 sieve was taken for the sieve analysis.
- 2. The sample was sieved through a set of sieves i.e. 4, 10, 14, 20, 40, 60, 70, 200 no. sieves. The sieving was performed with mechanical sieve shaker for 5 to 10 minutes.
- 3. The stack of sieves were removed from sieve shaker and weight of material retained on each sieve was computed. The

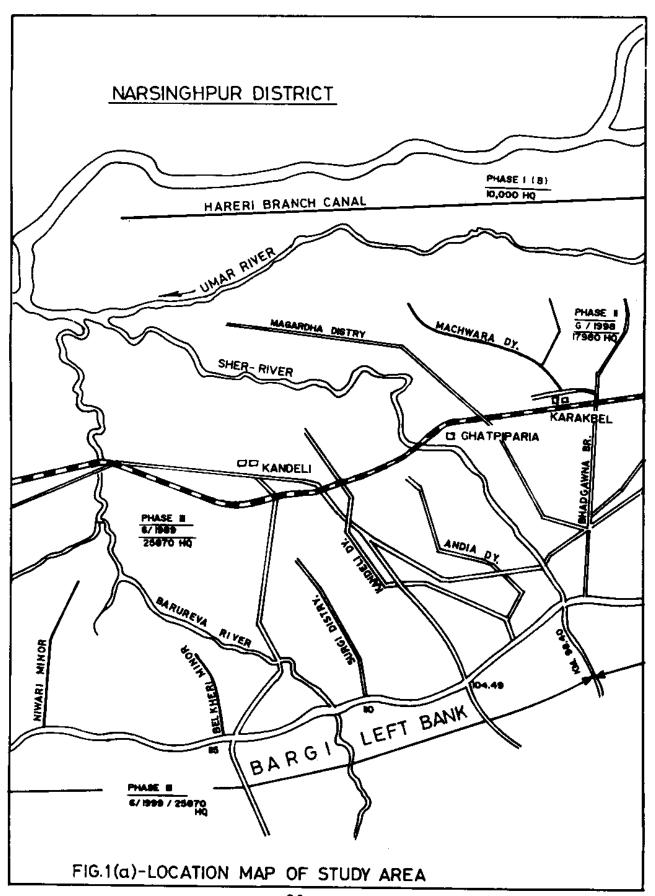


FIG. 2 LOCATION OF TEST POINTS AND GRIDS

- percentage of total soil sample retained on each sieve was also calculated.
- 4. The percentage of weight passing through each sieve was calculated. The calculation was started with 100 percent and subtracting the percentage retained on each sieve as a cumulative procedure as given by
  Percentage passing = Percentage arriving percentage

Percentage passing = Percentage arriving - percentage retained

- 5. A plot on semilogrithmic paper of grain size versus percent passing was plotted (fig 3 to 6).
- B. Hydrometer Analysis
- The soil group passing through the sieve no.200 was oven dried and 50 gm of the sample was taken for hydrometer analysis.
- The soil sample (50 gm) was soaked with 100 ml of sodium hexametaphosphate solution for 24 hours.
- 3. All the contents were transferred into the mixer cup and suitable quantity of water was added. The mixing was carried out for 2 to 3 minutes. The mixture was then carefully transferred to the sedimentation cylinder.
- 4. The cylinder of soil suspension was covered with rubber stopper and the suspension was carefully agitated for one minute. The jar was placed on table and the cap was removed. The hydrometer and thermometer were then inserted in the controlled jar and readings of hydrometer were taken out 0.5, 1,2,4,8,15,30,60,120, 240 and 1440 minutes.
- 5. The temperature of suspension was also recorded to the accuracy of  $1^{\circ}$ C for each hydrometer reading.
- 6. Between the hydrometer readings, hydrometer was placed in

another jar containing 100 ml of dispersive agent and 900ml of water and placed at the same temperature as other jar. The reading of hydrometer was calculated to find out the hydrometer corrections.

- 7. The diameter and the percentage finer with respect to each elapsed time interval was calculated. The grain size was plotted against percentage finer on semi log paper (fig 3 to 6).
- 3.3 Procedure for measurement of field saturated hydraulic conductivity
- 1. A well hole was prepared at the site with the help of soil auger and sizing auger. The hole was dug with the help of soil auger to a depth 15 cm less than the desired for final well depth. The 15 cm was dug with the help of sizing auger to produce a debris free well hole of uniform geometry of dia 6 cm and bottom flat.
- 2. Tripod was centered over the well hole and permeameter was lowered so that the support tube entered into the well hole.
- After the permeameter is placed, it is filled with water.
   Verification was then made for ensuring that both the reservoirs were connected.
- 4. The air inlet tip was slowly raised to establish the 5 cm well head height (H1).
- 5. The rate of fall of the water in the reservoir was measured at a regular time interval. The difference in readings at consecutive time interval divided by the time interval gave the rate of fall of water level, R1 in the reservoir.
- 6. The rate of fall of water in the reservoir was continuously monitored until it was almost same for three consecutive

- intervals. This rate of fall of water is called R1 and is defined as steady state rate of fall.
- 7. A 10 cm well head height (H2) was established and the rate of fall of water, R2, in the reservoir was obtained for stable value of R2.
- 8. The field saturated hydraulic conductivity (Kfs) was then calculated using following equation (9) (Tab 18).
  - 9. The matric flux potential  $(\phi_m)$  and alpha  $(\alpha)$  were calculated using equation (7) and (8) respectively (Tab 18).

### 4.0 Result And Discussion

The particle size analysis of the soil samples was carried out in the soil water laboratory of the Institute. The grain size distribution of the soil particles is given in Tab from 5 to From the particle size distribution curves the grain size distribution of the soil samples was found out. Using soil triangle the textural classification of the soil of study area was found and is given in Tab 17. The grain size distribution showed that in general sand content is varying from 5 to 18 %, except at N22 were sand content is as high as 47 %. Silt content is varying from 30 to 47 % and clay from 44 to 54 % except at N 22 were clay content is only 14 % (Tab 17). The particle size distribution curves are also prepared by plotting the grain size in millimeter and the percentage passing on a log normal scale for all the soil samples these graphs are illustrated in Figs from 3 to 6.

The saturated hydraulic conductivity of the soil samples measured by Guelph permeameter is given in tab (18). The insitu values are varrying from 0.005 to 0.6 m/day. Although at locations N21 the hydraulic conductivity value is very high about 2.7 m/day, this value is justified as the from USDA, drainage manual based on the textural analysis showed that soil at N 21 is sandy loam. At location N 52 the hydraulic conductivity obtained by Guelph was very low about 5E-05 m/day.

The relationship between in situ saturated hydraulic conductivity and texture has also been given by Johnson (1963). The hydraulic conductivity values obtained by Johnson's graph is given in Tab 19. The matric flux potential as calculated by equation (7) varried from 0.08 to 0.002 m<sup>2</sup>/day (Tab 18). The slope

#### Tab: 5 GRAIN SIZE DISTRIBUTION TEST OF N11

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 10-20 CM SAMPLING DATE OF TESTING: MARCH 23 1995

WT DRY SOIL = 39.0GMS
WT WASHED THRU NO 200 SIEVE = 461.0GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

## SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE

SIEVE	SIEVE D	WT RET	
МО	MM	GMS	% FINER
# 4	4.750	7.2	98.6
# 10	2.000	6.0	97.4
# 14	1.140	8.0	95.8
# 20	.825	8.6	94.0
# 40	.425	4.4	93.1
# 60	.250	1.0	92.9
# 70	.212	. 9	92.8
# 200	.075	2.5	92.3
PAN		. 2	

SUM SOIL WEIGHTS = 39.0 (39.0 GMS)

PROJECT: HYDROMETER ANALYSIS OF N11

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 10-20 CM SAMPLING

DATE OF TESTING: MARCH 23,1995

ZERO CORRECTION = 3.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.53 WT OF DRY SOIL WS = 50.00GMS

### HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM	•	DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	23.	50.0	98.2	.0545	90.55
1.0	23.	49.0	96.2	.0389	88.66
2.0	23.	48.0	94.1	.0278	86.76
4.0	23.	46.0	90.0	.0200	82.96
8.0	23.	44.0	85.9	.0144	79.16
15.0	23.	42.0	81.7	.0107	75.37
30.0	24.	39.0	76.2	.0077	70.24
60.0	24.	36.0	70.0	.0056	64.55
120,0	25.	<b>3</b> 3.0	64.4	.0040	59.42
240.0	2 <b>5</b> .	30.0	59.0	.0029	54.39
1440.0	27.	24.0	47.4	.0012	43.66

#### Tab:6 GRAIN SIZE DISTRIBUTION TEST OF N12

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOIL DESCRIPTION: 5-35 CM SAMPLING DATE OF TESTING: MARCH 15, 1995

WT DRY SOIL = 37.5GMS
WT WASHED THRU NO 200 SIEVE = 462.5GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE

SIEVE	SIEVE D	WT RET	
NO	MM	GMS	% FINER
# 4	4.750	6.1	98.8
# 10	2.000	6.4	97.5
# 14	1.140	7.3	96.0
# 20	.825	8.4	94.4
# 40	.425	4.6	93.5
# 60	.250	1.1	93.2
# 70	.212	. 7	93.1
# 200	.075	2.5	92.6
PAN		. 2	

SUM SOIL WEIGHTS = 37.3 ( 37.5 GMS)

PROJECT: HYDROMETER ANALYSIS OF N12

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 5-35 CM SAMPLING DATE OF TESTING: MARCH 23,1995

ZERO CORRECTION = 3.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.56 WT OF DRY SOIL WS = 50.00GMS

HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	23.	51.0	99.5	.0534	92.06
1.0	23.	50.0	97.5	.0381	90.17
2.0	23.	48.0	93.4	.0275	86.39
4.0	23.	46.0	89.3	.0198	82.61
8.0	23.	44.0	85.2	.0143	78.83
15.0	23.	42.0	81.1	.0106	75.05
30.0	24.	38.0	73.6	.0077	68.05
60.0	24.	36.0	69.5	.0055	64.27
120.0	25.	33.0	64.0	.0040	59.17
240.0	25.	31.0	60.6	.0028	56.05
1440.0	27.	24.0	47.0	.0012	43.48

### Tab: 7 GRAIN SIZE DISTRIBUTION TEST OF N21

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 10-40 CM SAMPLING

DATE OF TESTING: MARCH 18 1995

WT DRY SOIL = 132.3GMS
WT WASHED THRU NO 200 SIEVE = 367.7GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

## SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE

SIEVE	SIEVE D	WT RET	
NO	MM	GMS	% FINER
# 4	4.750	29.4	94.1
# 10	2.000	13.4	91.4
# 14	1.140	22.9	86.9
# 20	.825	33.4	80.2
# 40	.425	17.4	76.7
# 60	.250	3.9	75.9
# 70	.212	2.6	75.4
# 200	.075	8.9	73.6
PAN		. 3	

SUM SOIL WEIGHTS = 132.2 ( 132.3 GMS)

PROJECT: HYDROMETER ANALYSIS OF N21

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 10-40 CM SAMPLING

DATE OF TESTING: MARCH 23,1995

ZERO CORRECTION = 3.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.53 WT OF DRY SOIL WS = 50.00GMS

### HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

			· · · · · · · · · · · · · · · · · · ·		
ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	23.	50.0	98.2	.0545	72.22
1.0	23.	49.0	96.2	.0389	70.71
2.0	23.	48.0	94.1	.0278	69.20
4.0	23.	46.0	90.0	.0200	66.17
8.0	23.	44.0	85.9	.0144	63.14
15.0	23.	41.0	79.7	.0108	58.60
30.0	24.	39.0	76.2	.0077	56.02
60.0	24.	36.0	70.0	.0056	51.48
120.0	25.	34.0	66.5	.0040	48.91
240.0	25.	32.0	63.1	.0028	46.41
1440.0	27.	26.0	51.5	.0012	37.85

Tab:8 GRAIN SIZE DISTRIBUTION TEST OF N22

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 80-120 CM SAMPLING

DATE OF TESTING: MARCH 18 1995

WT DRY SOIL = 258.1GMS
WT WASHED THRU NO 200 SIEVE = 241.9GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE SIEVE D WT RET % FINER GMS NO MM 15.4 96.9 4.750 # 4 1.9 1.7 96.6 2.000 # 10 96.2 # 14 1.140 .825 4.1 30.2 75.5 37.7 95.4 # 20 89.3 74.3 66.7 .425 # 40 .250 # 60 .212 # 70 87.8 49.2 .075 # 200 2.6 PAN

SUM SOIL WEIGHTS = 256.8 ( 258.1 GMS)

PROJECT: HYDROMETER ANALYSIS OF N22

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 80-120 CM SAMPLING

DATE OF TESTING: MARCH 24,1995

ZERO CORRECTION = 2.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.59 WT OF DRY SOIL WS = 50.00GMS

HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

**					
ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
.5	27.	49.0	99.4	.0516	48.09
1.0	27.	46.0	93.3	.0376	45.14
2.0	27.	41.0	83.2	.0278	40.24
4.0	27.	37.0	75.0	.0203	36.31
8.0	27.	33.0	66.9	.0148	32.39
15.0	27.	29.0	58.8	.0112	28.46
30.0	27.	25.0	50.7	.0081	24.53
60.0	27.	22.0	44.6	.0059	21.59
120.0	27.	19.0	38.5	.0042	18.65
240.0	27.	17.0	34.5	.0030	16.68
1440.0	27.	12.0	24.3	.0013	11.78

### Tab:9 GRAIN SIZE DISTRIBUTION TEST OF N31

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 10-40 CM SAMPLING

DATE OF TESTING: MARCH 15 1995

WT DRY SOIL = 44.7GMS
WT WASHED THRU NO 200 SIEVE = 455.5GMS
TOTAL WT OF SOIL FOR TEST = 500.2GMS

### SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE

SIEVE	SIEVE D	WT RET	
NO	MM	GMS	% FINER
# 4	4.750	8.6	98.3
# 10	2.000	5.1	97.3
# 14	1.140	3.7	96.5
# 20	.825	4.1	95.7
# 40	.425	3.6	95.0
# 60	.250	3.4	94.3
# 70	.212	3.2	93.7
# 200	.075	12.1	91.3
PAN		.8	

SUM SOIL WEIGHTS = 44.5 ( 44.7 GMS)

PROJECT: HYDROMETER ANALYSIS OF N31

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 10-40 CM SAMPLING

DATE OF TESTING: MARCH 24,1995

ZERO CORRECTION = 2.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.56 WT OF DRY SOIL WS = 50.00GMS

#### HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	*FINERT
. 5	27.	48.0	98.1	.0526	89.35
1.0	27.	47.0	96.0	.0376	87.49
2.0	27.	46.0	94.0	.0268	85.63
4.0	27.	44.0	89.9	.0193	81.90
8.0	27.	41.0	83.8	.C140	76.32
15.0	27.	39.0	79.7	.0104	72.60
30.0	27.	38.0	77.6	.0074	70.74
60.0	27.	35.0	71.5	.0054	65.15
120.0	27.	33.0	67.4	.0039	61.43
240.0	27.	31.0	63.3	.0028	57.71
1440.0	27.	26.0	53.1	.0012	48.40

PROJECT: GRAIN SIZE DISTRIBUTION TEST OF N32

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 140-160 CM SAMPLING

DATE OF TESTING: MARCH 18 1995

WT DRY SOIL = 51.3GMS
WT WASHED THRU NO 200 SIEVE = 448.7GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE D WT RET SIEVE MM % FINER GMS NO 4.750 4.0 99.2 # 4 97.8 # 10 2.000 7.1 5.9 96.6 # 14 1.140 # 20 .825 10.3 94.5 7.4 .425 93.1 # 40 .250 # 60 3.5 92.4 .212 2.1 91.9 # 70 10.7 89.8 # 200 .075 . 2 PAN SUM SOIL WEIGHTS = 51.2 ( 51.3 GMS)

PROJECT: HYDROMETER ANALYSIS OF N32

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: 140-160 CM SAMPLING

DATE OF TESTING: MARCH 23,1995

ZERO CORRECTION = 3.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.57 WT OF DRY SOIL WS = 50.00GMS

HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
.5	23.	51.0	99.3	. 0532	89.08
1.0	23.	50.0	97.2	.0380	87.25
2.0	23.	48.0	93.2	.0274	83.60
4.0	23.	46.0	89.1	.0198	79.94
8.0	23.	44.0	85.0	.0143	76.28
15.0	23.	42.0	80.9	.0106	72.62
30.0	24.	39.0	75.4	.0076	67.68
60.0	24.	38.0	73.4	.0054	65.85
120.0	25.	35.0	67.9	.0039	60.91
240.0	25.	33.0	63.8	.0028	57.25
1440.0	27.	28.0	55.0	.0012	49.39

PROJECT: GRAIN SIZE DISTRIBUTION TEST OF N41

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 10-40 CM SAMPLING DATE OF TESTING: MARCH 18 1995

WT DRY SOIL = 98.2GMS
WT WASHED THRU NO 200 SIEVE = 401.8GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

#### SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE SIEVE D WT RET NO MM GMS % FINER # 4 4.750 5.6 98.9 # 10 2.000 5.7 97.8 # 14 1.140 5.7 96.6 .825 # 20 6.8 95.3 # 40 .425 13.3 92.6 # 60 .250 18.6 88.9 # 70 .212 11.7 86.5 # 200 .075 30.2 80.5 PAN . 6

SUM SOIL WEIGHTS = 98.1 ( 98.2 GMS)

PROJECT: HYDROMETER ANALYSIS OF N41

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE).

SOIL DESCRIPTION: 10-40 CM SAMPLING

DATE OF TESTING: MARCH 24,1995

ZERO CORRECTION = 2.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.50 WT OF DRY SOIL WS = 50.00GMS

## HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	27.	48.0	99.6	.0536	80.05
1.0	27.	47.0	97.5	.0383	78.38
2.0	27.	46.0	95.5	.0273	76.71
4.0	27.	44.0	91.3	.0197	73.37
8.0	27.	42.0	87.2	.0142	70.04
15.0	27.	40.0	83.0	.0105	66.70
30.0	27.	38.0	78.9	.0076	63.37
60.0	27.	36.0	74.7	.0054	60.03
120.0	27.	34.0	70.6	.0039	56.70
240.0	27.	32.0	66.4	.0028	53.36
1440.0	27.	28.0	58.1	.0012	46.69

PROJECT: GRAIN SIZE DISTRIBUTION TEST OF N51

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 20-50 CM SAMPLING

DATE OF TESTING: MARCH 15 1995

WT DRY SOIL = 32.8GMS
WT WASHED THRU NO 200 SIEVE = 467.2GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE

SIEVE	SIEVE D	WT RET	
NO	MM	GMS	% FINER
# 4	4.750	.0	100.0
# 10	2.000	. 5	99.9
# 14	1.140	.8	99.8
# 20	.825	3.0	99.2
# 40	. 425	8.3	97.5
# 60	.250	4.6	96.6
# 70	.212	2.5	96.1
# 200	.075	12.0	93.7
PAN		. 9	

SUM SOIL WEIGHTS = 32.5 ( 32.8 GMS)

PROJECT: HYDROMETER ANALYSIS OF N51

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOIL DESCRIPTION: 20-50 CM SAMPLING

DATE OF TESTING: MARCH 23,1995

ZERO CORRECTION = 3.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.57 WT OF DRY SOIL WS = 50.00GMS

HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	23.	51.0	99.3	.0532	92.76
1.0	23.	48.0	93.2	.0388	87.04
2.0	23.	46.0	89.1	.0280	83.23
4.0	23.	44.0	85.0	.0202	79.42
8.0	23.	41.0	78.9	.0146	73.71
15.0	23.	38.0	72.8	.0110	68.00
30.0	24.	36.0	69.3	.0078	64.76
60.0	24.	34.0	65.2	.0056	60.95
120.0	25.	32.0	61.8	.0040	57.71
240.0	25.	31.0	59.7	.0028	55.81
1440.0	27.	26.0	51.0	.0012	47.62

PAN

PROJECT: GRAIN SIZE DISTRIBUTION TEST OF N52

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: 100-140 CM SAMPLING

DATE OF TESTING: MARCH 18 1995

WT DRY SOIL = 34.3GMS
WT WASHED THRU NO 200 SIEVE = 465.6GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE SIEVE D WT RET NO GMS % FINER MM .0 100.0 4.750 # 4 # 10 2.000 .0 100.0 1.140 .0 # 14 100.0 99.9 # 20 .825 . 3 .425 .250 .212 .8 99.8 # 40 1.0 99.6 2.0 99.2 29.0 93.4 # 60 # 70 # 200 .075

SUM SOIL WEIGHTS = 34.1 ( 34.3 GMS)

PROJECT: HYDROMETER ANALYSIS OF N52

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

1.0

SOIL DESCRIPTION: 100-140 CM SAMPLING

DATE OF TESTING: MARCH 24,1995

ZERO CORRECTION = 2.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.60 WT OF DRY SOIL WS = 50.00GMS

#### HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	27.	49.0	99.1	.0514	92.34
1.0	27.	48.0	97.1	.0367	90.45
2.0	27.	44.0	89.0	.0270	82.92
4.0	27.	41.0	83.0	.0196	77.26
8.0	27.	37.0	74.9	.0143	69.72
15.0	27.	34.0	68.8	.0107	64.07
30.0	27.	32.0	64.8	.0077	60.30
60.0	27.	30.0	60.7	.0055	56.53
120.0	27.	28.0	56.7	.0040	52.76
240.0	27.	26.0	52.6	.0028	49.00
1440.0	27.	22.0	44.5	.0012	41.46

### Tab:14 GRAIN SIZE DISTRIBUTION TEST OF NAYAPURA

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: YELLOWIS HBROWN SOIL

DATE OF TESTING: NOVEMBE 22,1994

WT DRY SOIL = 60.5GMS
WT WASHED THRU NO 200 SIEVE = 439.5GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

#### SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE SIEVE D WT RET NO MM % FINER GMS # 4 4.750 2.9 99.4 # 10 2.000 2.3 99.0 # 14 1.140 6.6 97.6 # 20 .825 11.3 95.4 # 40 .425 13.2 92.8

# 60 .250 6.1 91.5 # 70 .212 4.0 90.7 # 200 .075 13.1 88.1

PAN .9

SUM SOIL WEIGHTS = 60.3 ( 60.5 GMS)

PROJECT: HYDROMETER ANALYSIS OF NAYAPURA

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: YELLOWISH BROWN SOIL

DATE OF TESTING: DECEMBER 8,1994

ZERO CORRECTION = 4.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.50 WT OF DRY SOIL WS = 50.00GMS

## HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	20.	53.0	101.7	.0552	89.38
1.0	20.	51.0	97.5	.0399	85.73
2.0	20.	49.0	93.4	.0288	82.09
4.0	20.	47.5	90.3	.0207	79.35
8.0	20.	44.0	83.0	.0151	72.97
15.0	20.	41.0	76.8	.0113	67.49
30.0	20.	39.0	72.6	.0082	63.84
60.0	20.	36.5	67.4	.0059	59.28
120.0	20.	34.0	62.3	.0042	54.72
240.0	20.	31.0	56.0	.0031	49.25
1440.0	20.	23.0	39.4	.0013	34.66

### Tab: 15 GRAIN SIZE DISTRIBUTION TEST OF RICHA

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: YELLOWIS HBROWN SOIL

DATE OF TESTING: NOVEMBE 22,1994

WT DRY SOIL = 33.8GMS
WT WASHED THRU NO 200 SIEVE = 466.2GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE D WT RET

		" T TILL T	
NO	MM	GMS	% FINEŔ
# 4	4.750	. 8	99.8
# 10	2.000	1.1	99.6
# 14	1.140	1.9	99.3
# 20	.825	4.4	98.4
# 40	.425	6.9	97.0
# 60	.250	3.6	96.3
# 70	.212	2.5	95.8
# 200	.075	11.8	93.4
PAN		.8	

SUM SOIL WEIGHTS = 33.6 ( 33.8 GMS)

PROJECT: HYDROMETER ANALYSIS OF RICHA

LOC OF PROJECT: SOIL WATER LABORATORY(N.I.H. ROORKEE)

SOIL DESCRIPTION: YELLOWISH BROWN SOIL

DATE OF TESTING: DECEMBER 8,1994

ZERO CORRECTION = 4.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.63 WT OF DRY SOIL WS = 50.00GMS

HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	%FINERT
. 5	20.	53.0	98.4	.0530	91.79
1.0	20.	51.0	94.4	.0383	88.04
2.0	20.	49.0	90.4	.0276	84.30
4.0	20.	47.0	86.4	.0199	80.55
8.0	20.	43.0	78.4	.0146	73.06
15.0	20.	40.5	73.3	.0109	68.37
30.0	20.	39.0	70.3	.0078	65.56
60.0	20.	36.5	65.3	.0056	60.88
120.0	20.	35.0	62.3	.0040	58.07
240.0	20.	33.0	58.3	.0029	54.32
1440.0	20.	28.0	48.2	.0012	44.96

Tab:16 GRAIN SIZE DISTRIBUTION TEST OF PALA (N55)

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

SOILDESCRIPTION: YELLOWIS HBROWN SOIL

DATE OF TESTING: NOVEMBER 22,1994

WT DRY SOIL = 288.6GMS
WT WASHED THRU NO 200 SIEVE = 211.4GMS
TOTAL WT OF SOIL FOR TEST = 500.0GMS

#### SIEVE ANALYSIS FOR GRAIN DISTRIBUTION CURVE SIEVE D WT RET NO % FINER MM GMS .0 # 4 4.750 100.0 # 10 2.000 . 2 100.0 # 14 1.140 . 1 99.9 .825 # 20 1.6 99.6 .425 # 40 43.8 90.8 .250 # 60 73.1 89.0 # 70 .212 50.5 63.0 # 200 .075 102.6 42.4

SUM SOIL WEIGHTS = 288.4 ( 288.6 GMS)

PROJECT: HYDROMETER ANALYSIS OF PALA

LOC OF PROJECT: SOIL WATER LABORATORY (N.I.H. ROORKEE)

. 6

SOIL DESCRIPTION: YELLOWISH BROWN SOIL

DATE OF TESTING: DECEMBER 8,1994

PAN

ZERO CORRECTION = 4.0 MENISCUS CORRECTION = 1.0 SPECIFIC GRAVITY = 2.61 WT OF DRY SOIL WS = 50.00GMS

### HYDROMETER ANALYSIS FOR GRAIN DISTRIBUTION CURVE

ELAP T	TEMP	HYDROM		DIAM	
MIN	DEG C	READING	% FINER	MM	XFINERT
. 5	20.	53.0	98.9	.0533	41.82
1.0	20.	49.0	90.8	.0393	38.41
2.0	20.	47.0	86.8	.0284	36.70
4.0	20.	44.0	80.7	.0206	34.14
8.0	20.	39.0	70.7	.0152	29.87
15.0	20.	35.5	63.6	.0115	26.89
30.0	20.	32.0	56.5	.0083	23.90
60.0	20.	30.0	52.5	.0060	22.19
120.0	20.	26.0	44.4	.0043	18.78
240.0	20.	25.0	42.4	.0031	17.92
1440.0	20.	21.0	34.3	.0013	14.51

Tab:17-Textural classification of Soil of Study Area

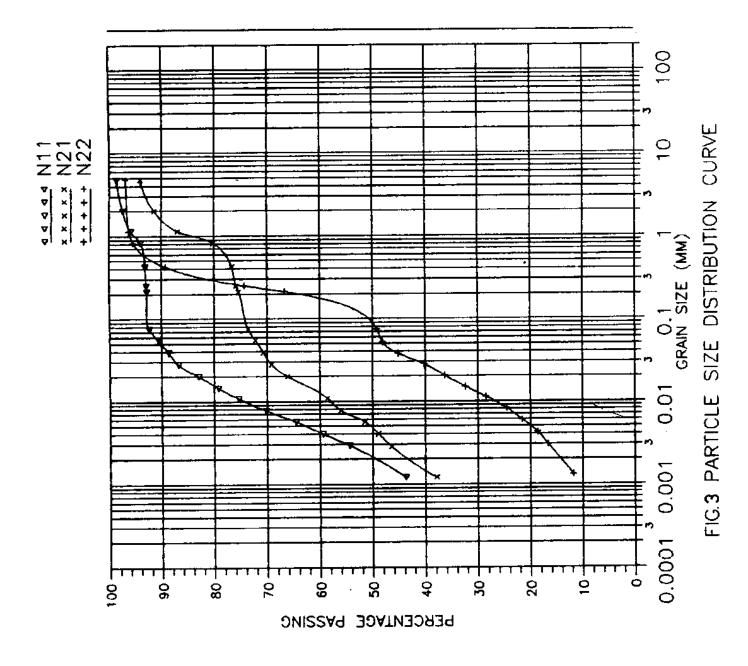
S. C. S.	i i	1	Location	Depth (cm)	Gravel *	Sand	Silt	Clay *	Textural Classifi- cation (USDA)
<del>.</del>	Z	11	Kheri	10-20	2.60	5.65	42.25	49.5	Silty clay
8.	z	12	Kheri	5-35	2.50	5.00	40.00	52.25	Silty clay
ო	z	21	Rampipariya	10-40	8.60	17.90	30.00	43.50	Silty clay
4.	z	22	Rampipariya	80-120	3.40	47.35	35.00	14.25	Sandy loam
, ,	z	31	Jallapur	10-40	2.70	7.05	36.25	54.00	Silty clay
. 9	Z	32	32 Jallapur	140-160	2.20	8.30	35.50	54.00	Silty clay
7.	z	41	Bhut pipariya	10-40	2.20	17.55	29.75	50.50	Silty clay
<b>œ</b>	z	51	Bahoripar	20-50	0.10	6.65	39.50	53.75	Silty clay
6	z	52	Bahoripar	100-140	0.00	7.00	47.25	45.75	Silty clay
10.	z	53	Noyapura	0-20	1.0	10.26	46.24	42.50	Silty clay
11.	z	54	Richa	0-20	0.40	7.0	42.10	50.5	Silty clay
12.	2	55	Pala	0-20	0.0	57.89	25.11	17.0	Sandy loam

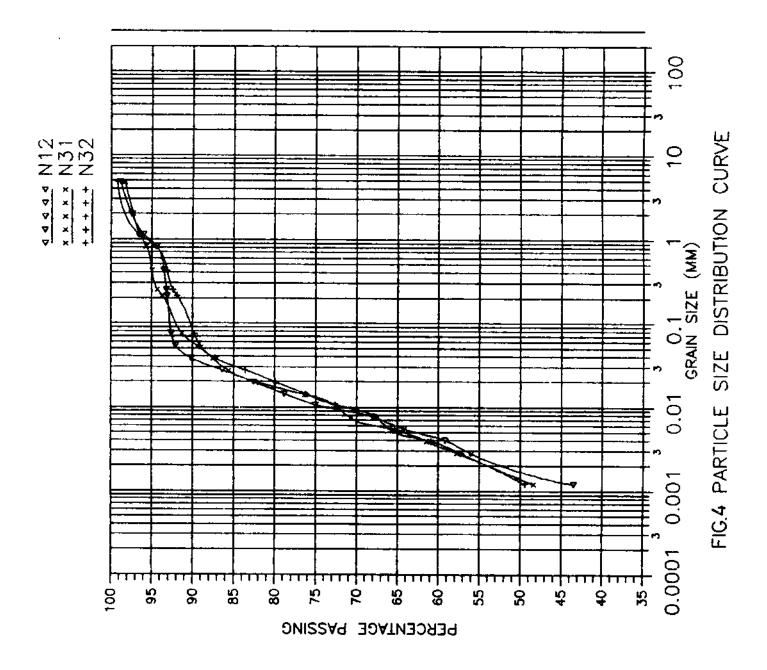
Tab.18: Saturated hydraulic conductivity Matric flux Potential and Alpha values for soil of Narsingpur

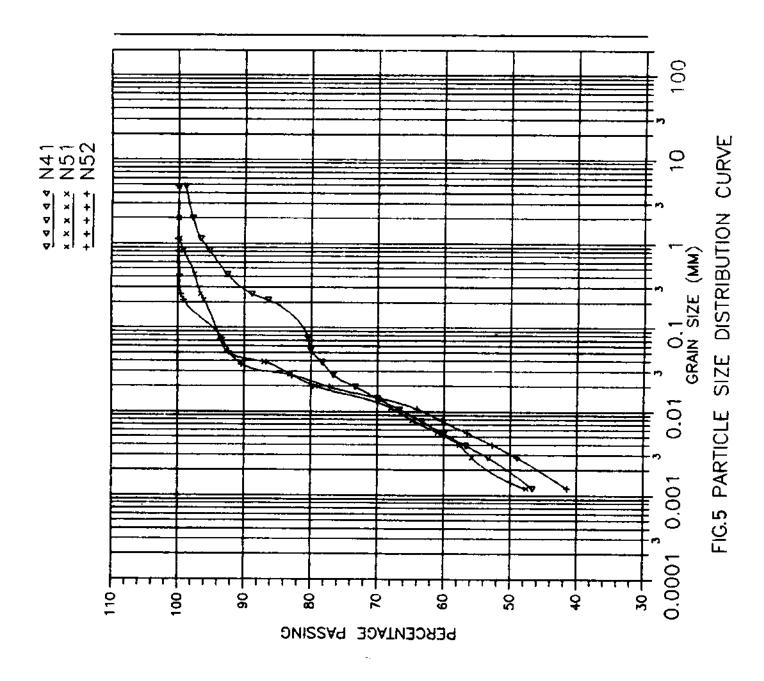
Sl	location	Saturated K	–	Alpha
no.	of test	(m/day)	Potential (m2/day)	(per m)
1.	N11	1.556790E-01	1.613628E-03	96,477620
2.	N12	5.713650E-01	5.376608E-03	106.268700
3.	N21	2.731425	8.393817E-02	32.540910
4.	N22	4.146301E-01	1.936354E-02	21.412940
5.	N31	4.680002E-02	3.614527E-03	12.947760
6.	N41	5.39 <b>2889E-0</b> 2	2.388170E-03	22.581680
7.	N52	5.780187E-05	1.613628E-03	3.582106E-02

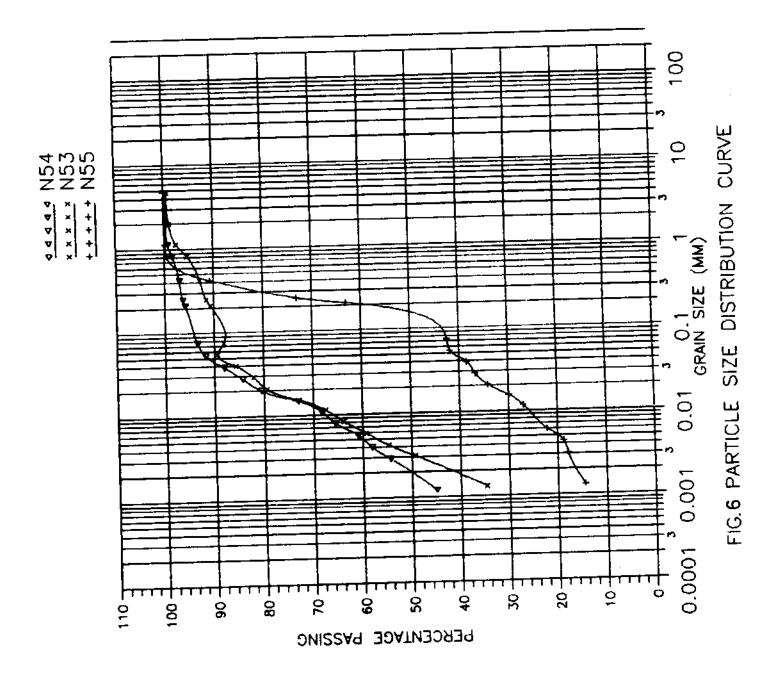
Tab 19 Saturated hydraulic conductivity by Guelph, Johnsons graph

Sl	location	Saturated	Conductivity	
no.	of test	Guelph (m/day)	Johnsons Graph (m/day)	
1.	N11	1.556790E-01	2.2E-03	
2.	N12	5.713650E-01	1E-03	
3.	N21	2.731425	0.02	
4.	N22	4.146301E-01	8E-04	
5.	N23	-	8E-04	
6.	N31	4.680002E-02	3E-03	
7.	N41	5.392889E-02	1E-03	
8.	N51	-	2E-03	
9.	N52	5.780187E-05	2E-03	
10.	N53	_	4E-04	
11.	N 5 4	_	8E-04	
12	N55	_	0.04	









Tab. 20 SOIL-MOISTURE CHARACTERISTICS CURVE FOR SILTY CLAY SOILS OF NARNADA SUBBASIN AT NARSINGPUR

# PRESSURE-MOISTURE CONTENTS FOR DIFFERENT SITES

	AVER	AGE MOIS	TURE CON	TENT AT	PRESSURE	(BAR)	
.10	.33	. 70	1.00	3.00	5.00	10.00	15.00
53.83	42.16	34.60	33.73	26.59	25.52	22.50	22.46
47.78	36.41	31.73	29.48	25.88	25.17	23.10	22.87
49.25	37.88	30.50	30.41	23.08	22.83	21.18	21.09
37.77	24.03	20.59	18.77	18.28	17.03	15.38	14.95
47.55	37.52	29.61	28.13	22.47	22.07	20.64	20.34
47.47	37.76	31.62	29.53	24.28	23.62	19.80	18.76
42.10	31.20	27.25	24.90	22.24	22.04	19.22	17.39
48.60	39.83	31.39	30.36	26.63	24.76	21.52	21.18
58.04	44.73	37.09	33.82	27.94	26.68	24.98	24.33



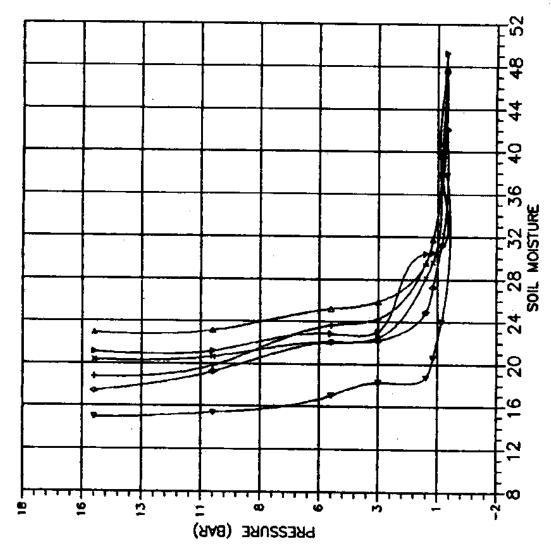


FIG. 7 SOIL MOISTURE CHARACTERISTIC CURVES

parameter alpha calculated using equation (8) varried from 13 to 106 per m (Tab 18), except for location N 52, where it was found very low i.e. 0.04 which is in accordance with rhe low hydraulic conductivity and matric flux.

The soil moisture characteristics obtained by using Pressure plate apparatus is given in tab 20. These values are plotted for 6 samples and soil moisture characteristic curves are obtained for the soil of the area (Fig 7).

## 5.0 Conclusions

From the tables it can be found out that the study area has predominant in clay. Soil survey organisation of Govt. of M. P.. has also showed that about 90 % of the area is containing mostly clay. Consequently the hydraulic conductivity and matric flux potential values are also low in accordance with the clayey texture of the soil.

# 6.0 Suggestions for further research

- The present study should be carried out on a smaller grid.
- Instead of the texture of the grid, the field experiments should be carried out on the basis of soil series given by Soil Survey office, Govt. of M.P..
- Analysis of ground water table data should also be carried out to find the general direction of ground water flow.

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